

**ATS 620**  
**Fall 2011**  
**Problem Set #6**  
**Due 21 October 2011 at 12 PM**

1. (10 points) In class, we derived an expression for  $\Delta E^*$  for heterogeneous ice nucleation on a planar substrate.

- a) Using that expression, plot values of  $\Delta E^*$  over the range of contact angles  $\theta = 0^\circ \rightarrow 180^\circ$  for elastic strain  $\epsilon = 0.0, 0.01, \text{ and } 0.02$ . Use the following values:

$$\begin{aligned}C &= 1.7 \times 10^{10} \text{ J m}^{-3} \\n_s &= 3.07 \times 10^{28} \text{ m}^{-3} \\ \sigma_{sv} &= 0.109 \text{ J m}^{-2} \\ T &= 263.15 \text{ K} \\ k &= 1.3087 \times 10^{-23} \text{ J K}^{-1} \\ e/e_s &= 1.1\end{aligned}$$

See Figure

- b) How do variations in contact angle and elastic strain affect the critical energy barrier? Physically, why should you expect this?

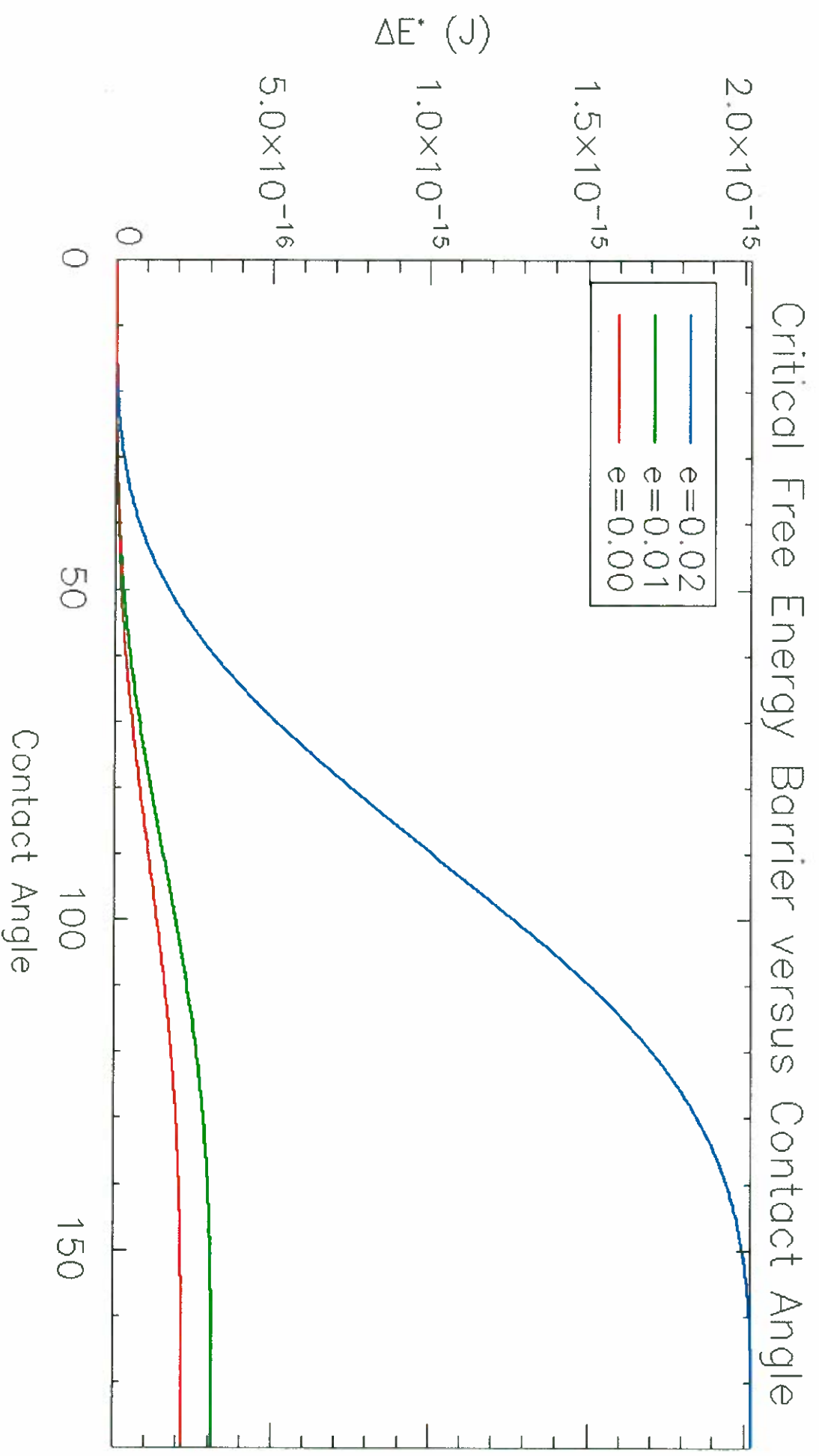
① As contact angle increases,  $\Delta E^*$  increases nonlinearly

② As strain increases,  $\Delta E^*$  increases nonlinearly

Explanations

① As contact angle increases, the dislocation between the ice lattice and the substrate acts to increase  $\sigma_{cs}$ , which increases  $f_c(m)$  and  $\therefore \Delta E^*$  increases.

② As elastic strain increases, the substrate looks less like ice. This puts "strain" on the bonds and acts to increase  $\mu_s$ . This increases  $\Delta E^*$



2. (10 points) Show the derivation of the equality

$$\Delta\mu = \mu_L - \mu_s = \boxed{RT \ln \frac{e_s}{e_i}} = \boxed{\frac{\Delta T_s}{T_0}}$$

Term 3 - Supersat  
Term

Term 4 - Supercool  
Term

Interpret the meaning of the third and fourth terms, particularly with respect to nucleation.

Term 3

Begin with Gibbs free energy,

$$dg = -SdT + vdp = d\mu$$

↓  
= 0 for constant T

$$d\mu = vdp$$

$$d\mu = d(\mu_L - \mu_s) = (v_L - v_s) dp$$

For Term 3, assume supersaturation is dominant and we have no supercooling

\* Unlike when we did this for equilibrium with water/vapor, we cannot assume  $v_L \gg v_s$  but we can still assume  $v_L \ll v_v$  and  $v_s \ll v_v$

$$d\mu = d((\mu_L - \mu_v) - (\mu_s - \mu_v))$$

use ideal gas,  $d\mu = d(\mu_L - \mu_v) - d(\mu_s - \mu_v) = (v_L - v_v) dp - (v_s - v_v) dp$

$$\hookrightarrow d(\mu_L - \mu_v) = -v_v dp = -\int_{e_s}^e RT d \ln e = -RT \ln \left( \frac{e}{e_s} \right)$$

$$d(\mu_s - \mu_v) = -v_v dp = -\int_{e_i}^e RT d \ln e = -RT \ln \left( \frac{e}{e_i} \right)$$

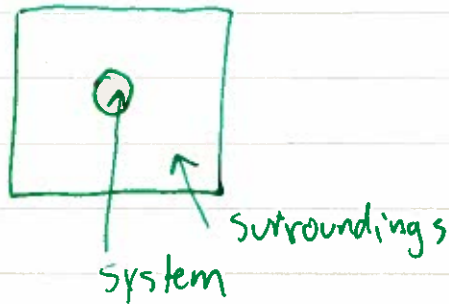
$$d\mu = d((\mu_L - \mu_v) - (\mu_s - \mu_v)) = d(\mu_L - \mu_s) = -RT \ln \left( \frac{e}{e_s} \right) - (-RT \ln \left( \frac{e}{e_i} \right))$$

$$d\mu = RT \left( \ln \frac{e}{e_i} - \ln \frac{e}{e_s} \right) = \boxed{RT \ln \left( \frac{e_s}{e_i} \right)} \text{ Term 3}$$

## Term 4

\* From First Law of Thermo  $\rightarrow Q_{in} = Q_{out}$

Consider a system with a supercooled drop in an environment



For a drop that is supercooled, let's assume the change in chemical potential is purely due to supercooling and not supersaturation, i.e.

$$\mu_L - \mu_S = \mu = -s dT + \underbrace{v dp}_0 = -s dT$$

$$d\mu_{LS} = -(s_L - s_S) dT$$

This is the change in entropy of the freezing supercooled drop

Going back to HW#1 P3, we need to calculate  $\Delta S$  of the Universe and break our irreversible process into reversible processes.

$$\begin{aligned} \text{For drop: } \Delta S &= \Delta S_{\text{warming}} + \Delta S_{\text{freezing}} + \Delta S_{\text{cooling}} \\ &= \int_{T_s}^{T_0} \frac{dQ_{\text{rev}}}{T} + \frac{L_f}{T} + \int_{T_0}^{T_s} \frac{dQ_{\text{rev}}}{T} \end{aligned}$$

But,  $\Delta S_{\text{univers}} = \Delta S_{\text{system}} + \Delta S_{\text{surroundings}} \quad \& \quad Q_{in} = Q_{out}$

We can now write our  $\Delta S$  for drop as:

$$\Delta S_{\text{system}} = \int_{T_s}^{T_0} \frac{dQ_{\text{in}}}{T} + \frac{L_f}{T} + \int_{T_0}^{T_s} \frac{dQ_{\text{out}}}{T}$$

+                      ↑                      ↓

Warming water                      Cooling of ice

$$\Delta S_{\text{surroundings}} = \int_{T_s}^{T_0} \frac{dQ_{\text{out}}}{T} + \int_{T_0}^{T_s} \frac{dQ_{\text{in}}}{T}$$

Using  $|Q_{\text{in}}| = |Q_{\text{out}}|$  or  $Q_{\text{in}} = -Q_{\text{out}}$ , we have

$$\Delta S_{\text{universe}} \Big|_L^S = \frac{L_f}{T_0} \quad ; \quad \text{but } \text{b/c we are going from liq to ice}$$

this must be negative.

Plugging this back into our  $\Delta \mu$  equation and remembering this process happened in a  $T_0$  to  $T_s$  range,

$$d\mu = -\Delta S dT$$

$$\int_{\mu_s}^{\mu_L} d\mu = \frac{L_f}{T_0} \int_{T_s}^{T_0} dT$$

$$\Delta \mu = \mu_L - \mu_s = L_f \frac{\Delta T_s}{T_0} \quad \text{where } T_s \equiv T_0 - T_s$$

3. (10 points) Problem 9.1 in Rogers and Yau:

An ice crystal in the form of a thin hexagonal plate grows by diffusion in an environment saturated with respect to water at a temperature of  $-4^\circ\text{C}$  and a pressure of 80 kPa. Determine the time required for it to grow to a diameter of 1 mm, starting from a mass of  $10^{-8}$  g. Take the capacity to be that of the circumscribing disk. Assume that the mass and the diameter of the plate are related by  $m = 1.9 \times 10^{-2} D^3$ , with mass in g and D in cm. Neglect ventilation effects. If diameter and fall speed are related by  $u = xD$ , where  $x = 520 \text{ s}^{-1}$ , determine the distance the crystal falls during the growth to 1 mm diameter.

$$\textcircled{1} \quad \frac{dm}{dt} = \frac{4\pi C(S_i - 1)}{\underbrace{\left(\frac{L_s}{kT} \left[\frac{L_s}{RvT} - 1\right]\right)}_A + \underbrace{\frac{RvT}{e_i(T)Dv}}_B} ; \quad m = 1.9 \times 10^{-2} D^3$$

$$\textcircled{2} \quad \frac{dm(t)}{dt} = (3)(1.9 \times 10^{-2}) D(t)^2 \frac{dD(t)}{dt}$$

Equate  $\textcircled{1}$  and  $\textcircled{2}$  and using  $C = \frac{D}{\pi r}$

$$\frac{dD}{dt} = \left( \frac{1}{D^2 (3)(1.9e^{-2})} \right) \left( \frac{4D(S_i - 1)}{A + B} \right)$$

$$D \frac{dD}{dt} = \frac{4}{(3)(1.9e^{-2})} \left( \frac{S_i - 1}{A + B} \right)$$

Integrate,  $\int_{D_i}^{D_f} D dD = \text{Bob} \int_{t_i}^{t_f} dt \Rightarrow \frac{D_f^2 - D_i^2}{2} = \text{Bob} \Delta t$

$$\Delta t = \frac{D_f^2 - D_i^2}{(2) \text{Bob}}$$

$$D_i = 8.07388 \times 10^{-5} \text{ m} ; \quad D_f = 1.0 \times 10^{-3} \text{ m}$$

(from m/D relation)

$$\Delta t = \frac{9.93481 \times 10^{-7} \text{ m}^2}{(2) (6.7360685 \times 10^9 \text{ m}^2 \text{ s}^{-1})} = \boxed{3624.0 \text{ seconds} = 55.77 \text{ min}}$$

$$\Delta H = v_0 T + \frac{1}{2} a T^2 ; \quad v_0 = (520 \text{ s}^{-1})(8.07388 \times 10^{-5} \text{ m}) = 0.04198 \text{ m s}^{-1}$$

$$a = 520 \frac{dD}{dt} = \frac{dV}{dt} = (520 \text{ s}^{-1}) \left( \frac{D_f - D_i}{\Delta t} \right) = 1.3e^{-4} \frac{\text{m}}{\text{s}}$$

$$T = 3624 \text{ sec}$$

$$\boxed{\Delta H = 1.018 \text{ km}}$$

4. (20 points) Using your programming language of choice, recreate the axes and isolines from slide 12 of Lecture 20. You do not need to turn in your code, however please turn in both your plot and the equations you used to create the plot.

Additional information: the vapor pressures over supercooled water and ice have remained uncertain. Extrapolations from values near the triple point are not accurate, even those in common use in the community. Recently Murphy and Koop (2005; paper posted on our class website) reviewed available thermodynamic data, uncovered problems with existing parameterizations, and presented new, more accurate parameterizations for supercooled water and ice vapor pressures. You should use these relationships in solving this problem (and in any other work you do in the below-0 C regime!).

See Plot

Equations used: From M&K (2005)

- Eqn (10)  $P_{iq}$
- Eqn (7)  $P_{ice}$

$$RH_{ice} = RH_w \frac{e_s(T)}{e_i(T)}$$

# Water and Ice Saturation Relationships

